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# The Gelasian Stage (Upper Pliocene): A new unit of the global standard chronostratigraphic scale

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*The Gelasian has been formally accepted as third (and uppermost) subdivision of the Pliocene Series, thus representing the Upper Pliocene. The Global Standard Stratotype-section and Point for the Gelasian is located in the Monte S. Nicola section (near Gela, Sicily, Italy).*

## Introduction

This short report announces the formal ratification of the Gelasian Stage as the uppermost subdivision of the Pliocene Series, which is now subdivided into three stages (Lower, Middle, and Upper). Furthermore, the Global Standard Stratotype-section and Point (GSSP) of the Gelasian is briefly presented and discussed. More information is contained in the paper by Rio et al. (1994), proposing the new chronostratigraphic unit, on which the formal procedure for ratification was based.

A postal ballot on the proposal by Rio et al. (1994) was forwarded to all voting members of the Subcommittee on Neogene Stratigraphy (SNS) in 1995 and accepted by a large majority (about 80%). Following the results of the postal ballot, a formal recommendation of SNS was submitted to the International Commission on Stratigraphy (ICS) in March 1996. Official acceptance by ICS and ratification by the Executive Committee of IUGS were achieved at the 30th International Geological Congress in Beijing (August 1996).

## Background and motivation

It is far beyond the scope of this report to review the development of Pliocene chronostratigraphy over the last century. The reader may find more information on the subject in Berggren (1971), Berggren & Van Couvering (1974), Cita (1975a) and Rio et al. (1991).

Even if numerous stages have been proposed over the years for subdividing the 3.5 Ma long Pliocene Series (Astian, Piacenzian, Zanclean, Fossanian, etc.), a two-folded subdivision into a Zanclean (Lower Pliocene) and Piacenzian (Upper Pliocene) Stages became well established in the eighties (e.g. Berggren et al., 1985; Haq and Van Eysinga, 1987).

Subsequently, a great effort was devoted to frame the chronostratigraphic units of the Pliocene within the increasingly more detailed biomagnetostratigraphy recently developed in the Mediterranean area (e.g. Rio et al., 1990; Sprovieri, 1992, 1993; Hilgen, 1991a, b, among many others). This approach led Rio et al. (1991, 1994) to demonstrate that none of the previously proposed Pliocene standard stages actually covered the interval of time between the top of the Piacenzian historical stratotype in Castell'Arquato (Northern Italy) and the Pliocene-Pleistocene boundary at Vrica (Southern Italy; Aguirre & Pasini, 1985).

Of course, this consideration alone does not imply that a new stage should be defined to represent the discovered gap. However, the top of the Piacenzian stratotype falls in a critical point of the evolution of Earth climatic system (i.e. close to the final build-up of the Northern Hemisphere Glaciation), which is characterized by plenty of signals (magnetostratigraphic, biostratigraphic, etc; see further on) with a worldwide correlation potential. Therefore, Rio et al. (1991, 1994) argued against the practice of extending the Piacenzian Stage up to the Pliocene-Pleistocene boundary and proposed the introduction of a new stage (initially "unnamed" in Rio et al., 1991), the Gelasian, in the Global Standard Chronostratigraphic Scale.

## The Gelasian Stage

As commented previously, a new chronostratigraphic unit is proposed to represent the interval of time above the Piacenzian historical stratotype and up to the already ratified GSSP of the Pleistocene Series.

The rank proposed is that of Stage which, according to the International Stratigraphic Guide (Salvador, 1994), is the "basic working unit of chronostratigraphy"; in addition "it is the smallest unit... that can be recognized at a global scale", a condition perfectly matched by the proposed Stage (see below).

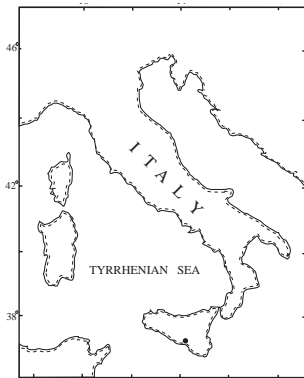
The name Gelasian is derived by the Greek name of the town of Gela (Southern Sicily), close to the boundary stratotype-section.

It is probably worth mentioning that the yellowish sandy sediments outcropping above the Piacenzian stratotype section in Castell'Arquato and below the local appearance of *Artica islandica*, that is a stratigraphic interval correlatable with the Gelasian, were ascribed to the Astian by Pareto (1865); however, the Astian, whose type is located in the Piedmont region (Northern Italy) (Mayer, 1868; Ferrero, 1971), has been abandoned, being represented by shallow water sandy facies, most probably Early or Middle Pliocene in age (Sampò et al., 1968).

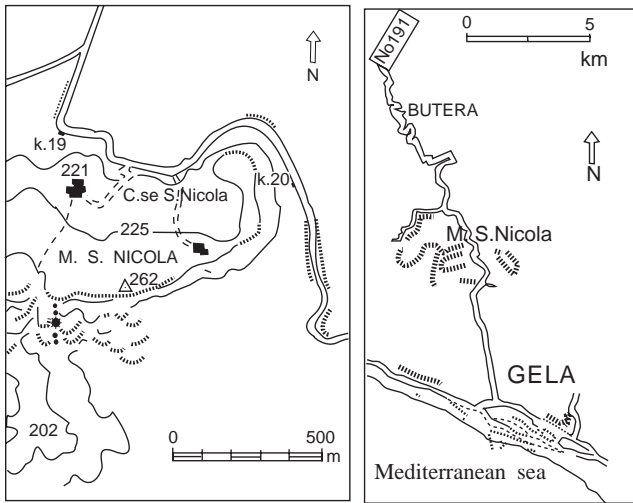
In the literature, some regional chronostratigraphic units (mainly regional stages) probably correlatable to (part of) the Gelasian Stage are found. For instance, the marine Nukumaruan Stage (Edwards, 1987) in New Zealand, the continental Tiglian stage (Zagwijn, 1974) in Northwestern Europe, and the continental "Middle Villafranchian" (Masini & Torre, 1990) in the Mediterranean.

## The GSSP of the Gelasian Stage

In this paragraph, we will briefly illustrate the characteristics of the Monte San Nicola section (Sicily), the definition of the "golden spike", and the various stratigraphic tools that provide an excellent potential for precise approximation of the boundary all over the world, in different geologic and stratigraphic settings. A more exhaustive account can be found in Rio et al. (1994).



**Figure 1**  
*Location of the Monte San Nicola section, the GSSP of the Gelasian Stage (modified after Rio et al., 1994).*



It is worth mentioning that another section was considered as candidate for the location of the base of the Gelasian Stage, that is the Singa section (Calabria, Southern Italy). Both sections have been thoroughly studied and well constrained in time (Channell et al., 1992; Sprovieri, 1992, 1993; Zijdeveld et al., 1991; Lourens et al., 1992), but outcropping and accessibility conditions are better at Monte San Nicola, that was thus preferred.

**The section**

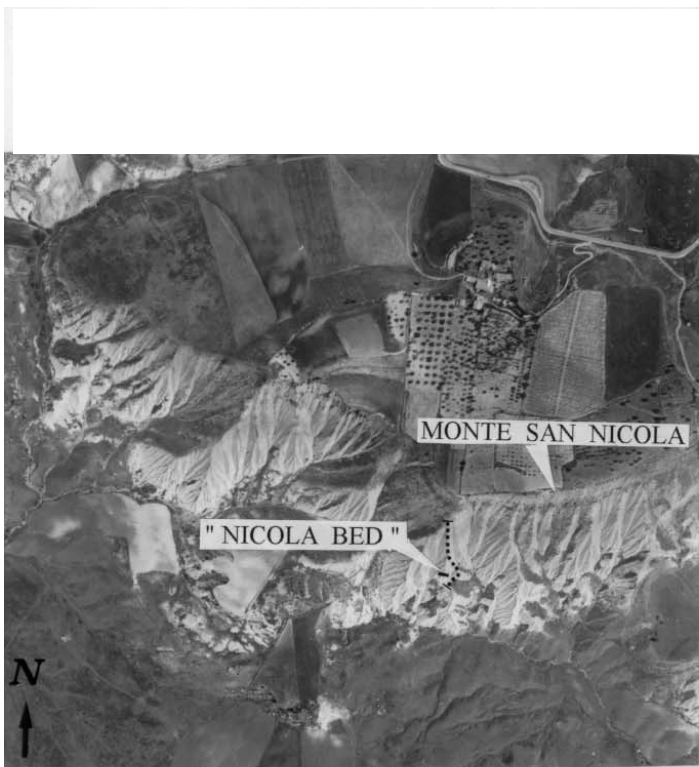
*Location*

The section outcrops in the badlands (“calanchi”) on the southern slope of the Monte San Nicola, some 10 km to the N-NW of the town of Gela (Caltanissetta province, Sicily) (Figure 1 and 2). The area is represented on the Carta Topografica d’Italia at 1:25000 Foglio 272, II NO (Ponte Olivo). The section can be reached, also by car, following a path which starts from the main road 191 between km 19 and 20 and leads to an old, little farm from which, crossing a small field by feet, the top of the section is easily reached.

*The stratigraphic succession*

The 161 m of the Plio-Pleistocene succession of the Monte San Nicola section were deposited on a piggy-back basin located on the so-called Gela Nappe (Argnani, 1987). They rest on sediments of the Numidian Flysch Formation (Oligocene-Lower Miocene; Gela Nappe). The Plio-Pleistocene strata dip 5-10° to the E and are unaffected by tectonic disturbances in the critical interval for the definition of the boundary. The depositional environment, at least for the boundary interval, is inferred to be a slope-basin setting with a water depth ranging from 500 to 1000 m (Rio et al., 1994, and references therein).

The lower 36 m are constituted by rhythmically bedded marls and limestones of the Trubi Formation (Cita and Gartner, 1973),



**Figure 2** Pictures of the Gelasian Stratotype section. In the lower right side, the arrow indicates the base of the marly layer overlying sapropel MPRS.

transitionally grading upward into the marly-silty Monte Narbone Formation. The latter is characterized by brownish-red laminated intervals ("sapropels" of Hilgen, 1991a, b and Lourens et al., 1992) and by brownish non-laminated manganeseiferous levels. Both are important for regional correlations; the "sapropels" have been demonstrated by Hilgen (1991a, b) to be an important tool for astronomical calibration of the Mediterranean sedimentary record.

The section was first described by Spaak (1983) who studied the planktonic foraminifers. Other studies are those by Bonaduce & Sprovieri (1984) (ostracodes), Driever (1988) (calcareous nannofossils), Channell et al. (1992) (biomagnetostratigraphy and paleoclimatology). Paleoenvironmental and cyclostratigraphic aspects of the section are dealt with in Sprovieri et al. (1986), Bertoldi et al. (1989), and Sprovieri (1992, 1993).

### Biomagnetostratigraphy, faunistic curve, sapropel stratigraphy

In the last years, the development of an integrated biostratigraphy (nannofossils and foraminifers), magnetostratigraphy and precise

lithostratigraphy/cyclostratigraphy of the rhythmically bedded sedimentary succession widely outcropping in Sicily and Calabria has resulted in an unprecedented time framework for the Mediterranean Pliocene and Pleistocene (among many others, Rio et al., 1990; Sprovieri, 1992, 1993; Hilgen, 1987, 1991a, b; Zijderveld et al., 1991; Lourens et al., 1992, 1996). This framework permits a precise bed-by-bed correlation of stratigraphic sections even tens or hundreds of kilometers apart, proving or disproving stratigraphic continuity of specific intervals. Besides, it allows a calibration of every single limestone-marl or marl-sapropel couplets of the entire sedimentary record to specific features in the curve of the quasi-periodic oscillations of the Earth's orbital parameters (Hilgen, 1991a, b; Lourens et al., 1996), leading to very precise (few kyrs) astronomical calibration of bioevents, magnetic boundaries and lithological record.

Concerning nannofossil biostratigraphy, the Monte San Nicola section is extended from Zone MNN14-15 to the top of Zone MNN19b (Rio et al., 1990), from Zone NN15 to NN19 (Martini, 1971), and from Zone CN11b to CN13a (Okada and Bukry, 1980). Concerning planktonic foraminifers, the section is extended from

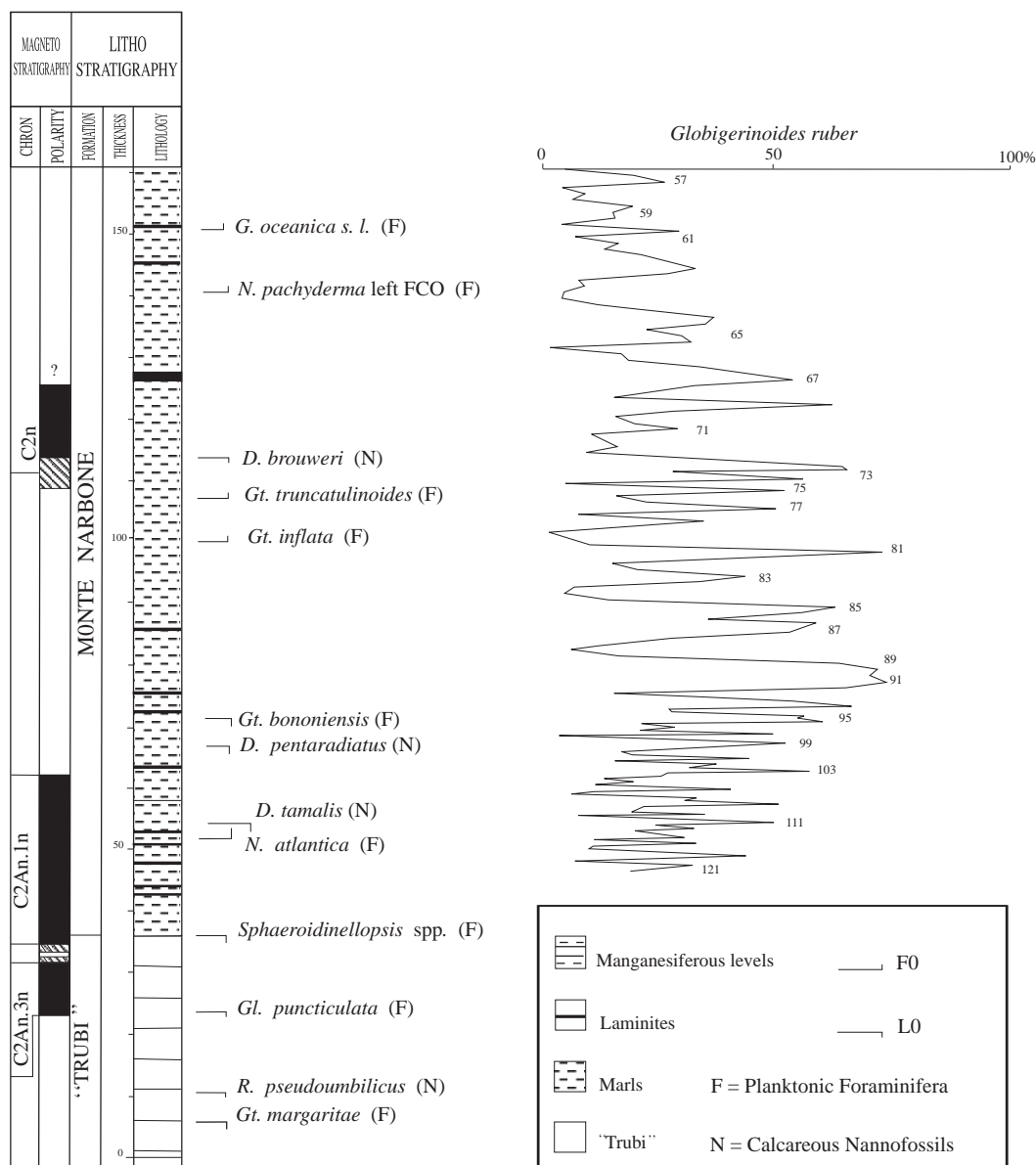


Figure 3 Summary of main magnetostratigraphic and biostratigraphic data available for the Monte San Nicola section. On the right side, the abundance fluctuations of *Globigerinoides ruber*, a surrogate for recognizing oxygen isotope stages (see text for further explanation) (after Rio et al., 1994).

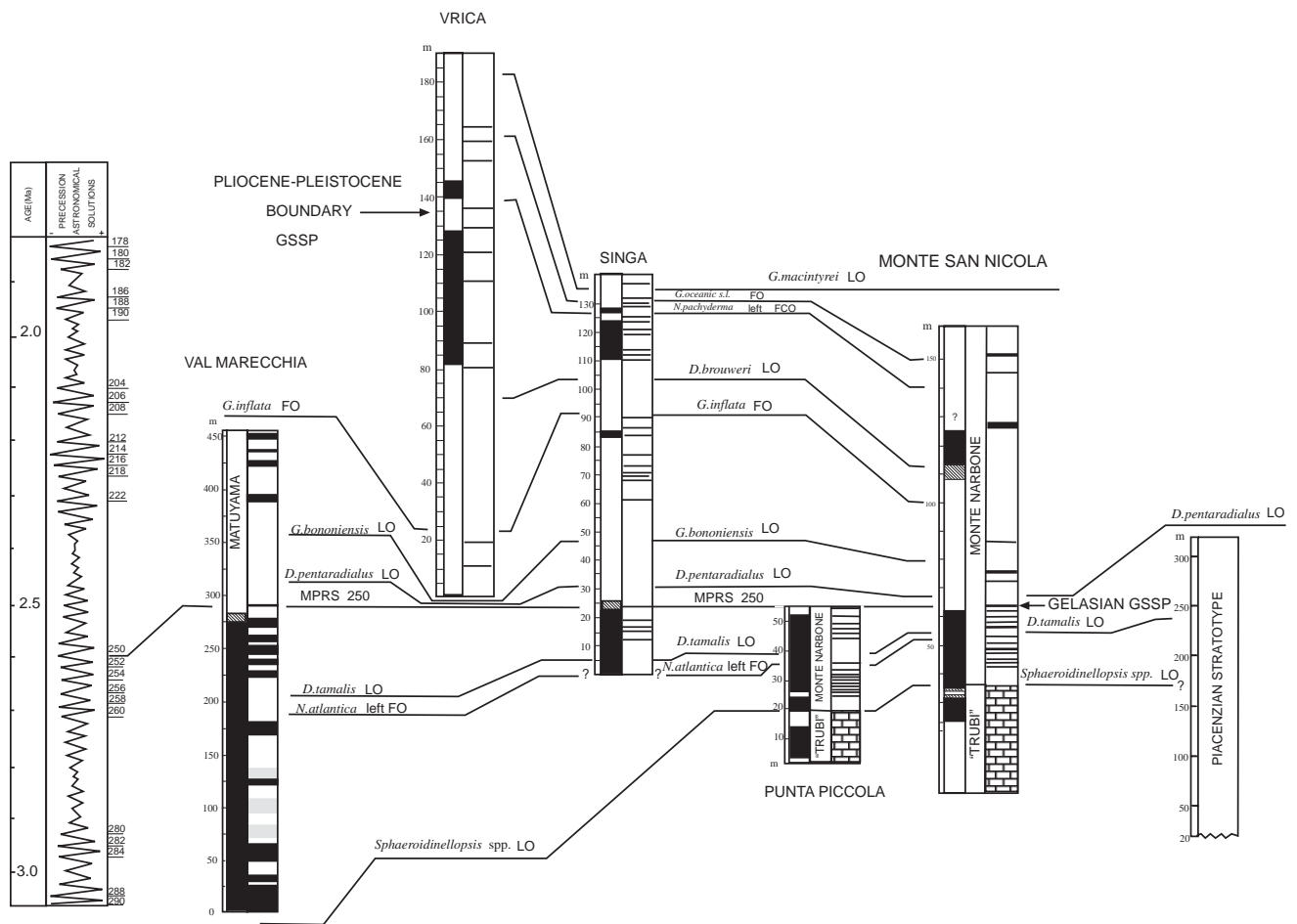


Figure 4 Correlation of different stratigraphic sections across the Piacenzian/Gelasian boundary from Northern Italy, to Calabria and Sicily, showing the high correlation potential of this chronostratigraphic boundary (after Rio et al., 1994).

Zone MPI3 to the *Globigerina cariacensis* Zone (Cita, 1973, 1975b, emended by Sprovieri, 1992), and from Zone N19 to N21 (Blow, 1969) (Figure 3).

From a magnetostratigraphic point of view, the section spans the interval from the upper part of the Gilbert Chron (C2Ar of Cande & Kent, 1992, 1995) to the Matuyama Chron above the Olduvai Subchron (C1r). The Gauss/Matuyama boundary is clearly detectable and precisely located (Channell et al., 1992) (Figure 3).

Abundance fluctuations of *Globigerinoides ruber*, a surrogate for recognizing oxygen isotopic stages (Sprovieri, 1993; see also Ruddimann et al., 1989, Lourens et al., 1992), allow the recognition in the Monte San Nicola section of all isotopic stages from 116 to 56 of Raymo et al. (1989) (Figure 3).

In the interval critical for the definition of the base of the Gelasian Stage, a cluster of six sapropels or manganeseiferous levels is recognized and correlated to coeval sections in Sicily (Punta Piccola), Calabria (Singa) and even Northern Italy (Val Marecchia). The single sapropels of this cluster (cluster "A" of Verhallen, 1987) have been coded by Hilgen (1991b) as Mediterranean Precession Related Sapropels (MPRS) and tightly linked, as already discussed, to the corresponding precessional cycle. One of this sapropel is selected as lithologic marker for the location of the "golden spike" (see below).

Based on all the evidences listed above, it can be confidently assumed that the Monte San Nicola section is remarkably complete in a thick stratigraphic interval straddling the selected boundary.

## The boundary ("golden spike")

### Definition and age

The base of the marly layer overlying sapropel MPRS 250, located at 62 m in the Monte San Nicola section, is the defined base of the Gelasian Stage (Middle Pliocene-Upper Pliocene boundary) (Figure 3 and Figure 4, right side).

The astrochronological age of sapropel MPRS 250 (mid-point), corresponding to precessional cycle 250 from the present, is 2.588 Ma (Lourens et al., 1996), which can be assumed as the age of the boundary.

### Correlation tools

The base of the Gelasian corresponds to isotopic stage 103 of Raymo et al. (1989), predating by about 60 kyrs the prominent cold stage 100, often referred to as final build-up of Northern Hemisphere Glaciation. This major climatic threshold is also witnessed by increase in ice rafted detritus in northern latitude oceanic sediments (Raymo et al., 1989), change in vegetation distribution patterns (Tiglian-Pretiglian boundary of Zagwijn, 1974; Arquatian phase of Lona, 1962), the beginning of loess sedimentation in China (Kukla & An, 1989) and migratory events in the continental vertebrate fauna (the elephant-Equus event in the Eurasian region; Lindsay et al., 1980; Azzaroli et al., 1988 and references therein).

The Gauss/Matuyama boundary, located about 1 m (20 kyrs) below the "golden spike", is a very good approximation of the base of the Gelasian outside the stratotype section.

From a biostratigraphic point of view, the extinction level (LO) of the nannofossils *Discoaster pentaradiatus* and *D. surculus* occur in most low- and mid-latitude areas close to isotopic stages 99, that is some 80 kyrs above the boundary. In the Mediterranean and North Atlantic the LO of the foraminifer *Globorotalia bononiensis* (or *G. punctulata* of some authors) is recorded in stage 96, i.e. some 140 kyrs above the boundary. The LO of the radiolaria *Stichocorys peregrina*, coincident with the base of the *Pterocanium prismaticum* Zone, approximates the Gauss-Matuyama boundary (Sanfilippo et al., 1985). The first occurrence of the diatom *Nitzschia jousseauei* in low-latitudes and the LO of *Denticulopsis kamtschatica* in the North Pacific mid- and high-latitudes approximate the Gauss-Matuyama boundary (Barron, 1985).

## Conclusion

The introduction of the Gelasian Stage (Upper Pliocene) and the formal definition of its base mark an important step towards the compilation of the Global Standard Chronostratigraphic Scale of the Neogene, which now include a three-folded Pliocene Series. In this issue, we also report on the acceptance and ratification of the base of the underlying Piacenzian Stage (Middle Pliocene).

The definition of the base of the Zanclean Stage (Lower Pliocene) is one of the major goals of SNS in the next few months: we hope to report on this important achievement soon.

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